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3	Investigating Impacts of Local Circulation on Coastal Ozone Pollution in the							
4	New York Metropolitan Area: Evidence from Multi-year Observations							
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13	Key Points:							
14	1. K-means clustering is applied to group local circulation and temperature diurnal cycles							
15	during the summertime of 2017-2019.							
16								
17	2. Hot summer day ozone would drop 9-10 ppb when sea breeze is delayed and intercepted							
18	by the dominant westerlies in ozone hot spots.							
19								
20	3. Early morning meridional wind direction would control sea breeze onset type and							
21	modulate ozone exceedances during extreme conditions.							

modulate ozone exceedances during extreme conditions.

22 Abstract

Elevated surface ozone levels are often detected in the New York (NY) metropolitan area during 23 summertime. Moreover, surface ozone in this region exhibits sharp spatial gradients and 24 25 distinctive diurnal cycles under the influence of complex boundary layer circulation induced by the intricate coastal geometry. This study examines how surface ozone is impacted by local 26 circulation spatially and temporally under different temperature scenarios (all summer days, hot 27 summer days, and extreme heat days) with the help of cluster-based meteorological conditions 28 29 during the summertime of 2017-2019. The most polluted days are found to be highly associated with hot sea breeze days with weak background flow. When sea breeze development in the New 30 31 York Bight is delayed and its penetration north is intercepted by the dominant westerlies during hot summer days, daily maximum 8-hour average ozone (DMA8) in some ozone hot spots of 32 33 New York City (NYC) and the south shore of Connecticut (CT) would typically drop 9-10 ppb under comparable temperature levels. The average regional decrease of DMA8 for NYC and 34 35 coastal CT is 6.7 and 8.3 ppb, respectively. Furthermore, we conclude that a change in early 36 morning meridional wind direction is the most critical meteorological characteristic in 37 controlling sea breeze onset type and helping modulate ozone exceedances in the region during 38 extreme heat days when ozone exceedances are expected to be very common. The conclusion is further demonstrated with two case studies during the Long Island Sound Tropospheric Ozone 39 Study (LISTOS) 2018 field campaign. 40

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42 Plain Language Summary

The New York metropolitan area is classified as a moderate 8-hour ozone nonattainment area by 43 the U.S. Environmental Protection Agency. Moreover, surface ozone in this region exhibits sharp 44 horizontal and vertical gradients and distinctive diurnal cycles under the influence of complex 45 boundary layer circulation. This study examined three summers of meteorological (wind and 46 47 temperature) and chemical (ozone and its precursors of nitrogen oxides and formaldehyde) 48 observations to investigate the impacts of local circulation on coastal ozone pollution in the area, with a focus on hot summer days when high concentrations of ozone are more likely. With the 49 projected extreme heat in the region increasing in intensity and duration throughout the 21st 50

- 51 century, meteorological characteristics most critical in modulating ozone exceedances in ozone
- 52 hot spots during extreme heat days are identified and analyzed with case studies during the Long
- 53 Island Sound Tropospheric Ozone Study (LISTOS) field campaign.

55 **1 Introduction**

The New York (NY) metropolitan area, including New York City (NYC), Long Island (LI), the 56 57 south shore of Connecticut (CT), and the northeast corner of New Jersey (NJ), is the most populous built-up area in the United States. Elevated surface ozone levels are often detected in 58 the region during summertime thanks to favorable meteorological conditions such as high 59 temperature and abundant local anthropogenic and natural ozone precursor emissions of nitrogen 60 oxides (NOx, consisting of NO and NO₂) and volatile organic compounds (VOCs). The region is 61 classified as a moderate 8-hour ozone nonattainment area (area that has a design value of 81 up 62 to but not including 93 ppb, US EPA, 2018) with the Ozone National Ambient Air Quality 63 Standards (NAAQS) level at 70 ppb. 64

Moreover, surface ozone in the NY metropolitan area exhibits sharp horizontal and vertical 65 66 gradients and distinctive diurnal cycles under the influence of complex boundary layer circulation, such as the commonly detected Atlantic sea breeze (sea breeze developed along the 67 south coast of LI) and sound breeze from Long Island Sound (LIS) induced by the intricate 68 coastal geometry throughout the region (Couillard et al., 2021; Han et al., 2022; McCabe & 69 70 Freedman, 2023; Novak & Colle, 2006; Zhang et al., 2020). Other large water bodies in the region, such as the Hudson River, Upper Bay, Lower Bay, and Newark Bay, also have the 71 72 potential to influence local circulation due to land/water temperature contrast. In addition, the 73 large-scale urbanization in the region alters the local circulation with its excessive heat flux and increased surface roughness. Urban heat island (UHI) generated by the heat contrast between the 74 urban center and surrounding area produces convergence flow into the urban center. The 75 76 convergent flow would accelerate the sea breeze front towards the city center (Freitas et al., 2007; 77 Hu et al., 2022). However, sea breeze front penetration over the city center would likely be more difficult and can be stalled for a few hours (Freitas at al., 2007; Hu et al., 2022) due to the 78 79 convergent flow and urban frictional retardation and would be slower overall than that over lessurbanized surrounding areas (Ferdiansyah et al., 2020; Han et al., 2022). These interactions have 80 been found to strongly affect urban air pollution (Bornstein & Thompson, 1981; Han et al., 2022; 81 82 Khiem et al., 2010; Ulpiani, 2021).

Impacts of local circulation on surface ozone have been studied across many coastal regions
worldwide, especially near or around urban centers where high ozone remains a threat to public

85 health. Many of them are case studies using observations (Banta et al., 2005; Couillard et al., 2021; Han et al., 2022; Zhang et al., 2022), numerical modeling (Ding et al., 2004; Torres-86 87 Vazquez et al., 2022) or both (Caicedo et al., 2019; Finardi et al., 2018; Liu et al., 2022). It has been found that large-scale flow direction, sea breeze characteristics, and vertical mixing all play 88 essential roles in controlling ozone horizontal and vertical distribution. Sea breeze characteristics 89 include location/penetration, timing, depth, and intensity/strength. Specifically, for the NY 90 region, Zhang et al. (2022) found that the north shore of LI experienced the highest ozone during 91 stagnant or sea breeze days, while the south shore high ozone always involved westerly or 92 southwesterly flow; Torres-Vazquez et al. (2022) found that ozone over CT is especially 93 sensitive to local wind features. Despite numerous case studies examining the influence of local 94 circulation on the coastal ozone problem in the NY metropolitan area, particularly during the 95 Long Island Sound Tropospheric Ozone Study (LISTOS) 2018 field campaign (e.g., Han et al., 96 2022; Torres-Vazquez et al., 2022; Zhang et al., 2022), a comprehensive investigation over an 97 extended period is warranted to better summarize local conditions and serve as a reference for 98 future case studies. 99

100 Long-term studies routinely employ group-based comparisons across diverse scenarios, such as sea breeze days and non-sea breeze days. Ngan & Byun (2011) and Li et al. (2020) examined the 101 102 ozone levels at various circulation conditions in Houston with gridded analysis products and insitu and remote sensing observations, respectively. The highest ozone is found to be associated 103 with sea breeze days in the Houston region (Li et al., 2020). Papanastasiou and Melas (2009) 104 studied the difference in observed ozone diurnal cycles during sea breeze versus non-sea breeze 105 106 days in the Volos area of Greece. They concluded that the presence of sea breeze will increase ozone concentrations, and the increase is more pronounced during cold months. Oh et al. (2006) 107 compared the observed diurnal variations in ozone and NO₂ during early and late sea breeze 108 onset days in the Busan metropolitan area in Korea and explored the mechanisms with case 109 studies using numerical modeling. It is found that offshore synoptic flows delay the onset of sea 110 breeze and help preserve urban emission above the sea from early morning, which will be 111 recirculated back with sea breeze, leading to a significant enhancement of maximum ozone by a 112 factor of 1.5 compared with early sea breeze onset condition. 113

However, these group-based comparison studies only focused on specific aspects of local 114 circulation and overlooked the importance of temperature to local ozone levels, despite the well-115 established strong correlation between surface ozone levels and temperature (Coates et al., 2016; 116 Olszyna et al., 1997; Porter & Heald, 2019). Higher temperatures often indicate clear sky 117 conditions with higher levels of solar radiation reaching the lower atmosphere. Together, they 118 favor higher surface ozone concentration by 1) accelerating photochemical reaction rates and 2) 119 boosting VOCs and NOx emissions from both natural and anthropogenic sources through various 120 processes, such as accelerated biogenic emission of VOCs and increased NOx emissions from 121 higher energy demand for air conditioning (Coates et al., 2016; Guenther et al., 1995; Porter & 122 Heald, 2019). In addition, higher temperature facilitates the growth of a deeper planetary 123 boundary layer (PBL) due to increased convection driven by surface heating. However, the 124 impacts of PBL height on surface ozone are often coupled with influences from other physical 125 and chemical factors, including the absolute height of PBL and ozone vertical profile (Haman et 126 al., 2014; Zhang et al., 2023). The overall effect of PBL height on surface ozone is less certain. 127 In general, the highest surface ozone most likely occurs when PBL height is moderate (between 128 129 1-2 km for Beijing) and decreases as PBL becomes higher (dilution) or lower (decreased photochemistry because of the availability of sunlight) (Zhang et al., 2023). 130

With the addition of temperature as a contributing factor, this study adapted the k-means clustering technique used in Li et al. (2020) to cluster NYC wind conditions and diurnal temperature profiles. It comprehensively investigated the impacts of local circulation on coastal ozone pollution in the area using three summers of meteorological and chemical observations. Specific questions we will answer in this study are:

What is the general relationship between surface ozone and meteorological conditions in the
 region during summertime?

What are the temporal and spatial characteristics of ozone during hot summer days, when
elevated ozone levels are more prevalent, and how does local circulation contribute to these
patterns?

141 3. Can local circulation help modulate ozone exceedances during extreme heat events, and if142 so, how?

The contents are organized as follows. Details of data and methodologies are described in 143 Section 2. Section 3.1 addresses the first question by presenting the cluster results and examining 144 the overall impact of local circulation and temperature scenarios on daily maximum 8-hour 145 average ozone (DMA8) spatial distribution in the New York metropolitan area during the 146 summertime of 2017-2019. Then, Section 3.2 focuses on hot summer days, during which ozone 147 levels are more elevated and harmful to human health and other living beings, and summarizes 148 how background wind conditions and sea breeze impact ozone diurnal cycles and spatial 149 distributions, answering the second question. Lastly, critical meteorological characteristics that 150 can help modulate ozone exceedances during extreme hot sea breeze days in ozone hot spots of 151 NYC and coastal CT are identified and illustrated with representative case studies during 152 LISTOS in Section 3.3 (the third question). A comprehensive table detailing all related 153 154 meteorological characteristics and ozone levels during hot sea breeze days is also included to serve as a reference for future case studies. 155

156 **2 Data and Methods**

157 **2.1 Data**

Multi-platform meteorological and chemical observations during the summertime (June-July-August, JJA) of 2017-2019 were collected to help answer the three scientific questions above. All surface and lidar observations with temporal resolution higher than an hour have been averaged to hourly data for consistency and more straightforward interpretation. The hour for all hourly data indicates the end of the averaging window in EST (UTC - 05:00, EST is used throughout this study), e.g., 11:01 to 12:00 is averaged for the 12:00 condition for data measured at 1 min frequency.

Hourly surface ozone observations and hourly NOx and meteorological observations at available 165 co-located sites were directly obtained from the pre-generated annual files from the US 166 (EPA) 167 Environmental Protection Agency Air Quality System website (AQS, 168 https://aqs.epa.gov/aqsweb/airdata/download_files.html). Since they are reported at the sample starting time, all AQS observations are shifted one hour to be consistent. Ozone, NOx, and NO₂ 169 observations used in this study are measured by either a Federal Reference Method (FRM) or 170 171 a Federal Equivalent Method (FEM) as defined in the Code of Federal Regulations for Ambient Air Monitoring Reference and Equivalent Methods (https://www.ecfr.gov/current/title-40/chapter-I/subchapter-C/part-53). Locations of these air quality monitoring sites are marked with red triangles in Figure 1. Specifically, site 360810124 at Queens, NY, and 090019003 at Westport, CT, both with NO₂ observations, are annotated in red on the map. Both sites have been utilized to study how local circulation modulates ozone exceedances during extreme hot sea breeze days in the highly polluted NYC and coastal CT regions.

Additional hourly surface meteorological observations were assembled from the New York State 178 Mesonet (NYSM, https://www2.nysmesonet.org/) (Brotzge et al., 2020), the Automated Surface 179 Observing Systems (ASOS, https://www.ncei.noaa.gov/products/land-based-station), and the 180 181 National Data Buoy Center (NDBC, https://www.ndbc.noaa.gov/). Their locations, including those of those AQS sites with meteorological observations, are shown with black dots in Figure 1. 182 183 NYSM performs both automated and manual quality control to flag out erroneous data. Pregenerated hourly NYSM data was obtained directly by data request at the official website. One-184 185 minute observations from ASOS were downloaded and averaged to the end of the hour. Buoy data is reported at various temporal resolutions from several minutes to one hour at the end-of-186 acquisition time. The same averaging was conducted for those observations with higher temporal 187 resolutions. Specifically, hourly data is reported at two buoy sites (44025 and 44064) used to 188 189 represent air temperature in the New York Bight at hh:50 and shifted ten minutes to (hh+1):00.

Among these, the newly established NYSM delivers quality-controlled hourly average 190 191 meteorological fields with extraordinary temporal data coverage. For example, site QUEE at 192 Queens, NY, experienced only two missing days for temperature and one missing day for surface wind during the three summers we studied. Thus, eight NYSM surface sites (thick gray circles in 193 Figure 1), located in NYC and LI, in addition to those two buoy sites, 44025 and 44065 (thick 194 195 gray circles in the ocean in Figure 1), both located south of LI, were adopted to calculate 196 regional meteorological conditions over land and sea, respectively. Among these eight NYSM surface sites, Queens, NY (QUEE) in NYC and Wantagh, NY (WANT) in LI were used to 197 198 cluster local weather conditions and annotated in black on the map.

NYSM wind Doppler lidar (WDL, Leosphere scanning Doppler Windcube 100S) collects data
using the Doppler Beam Swinging (DBS) scan mode, which consists of five scans in four
cardinal directions (north, east, south, and west) and nadir with a cycle of 20 seconds (Shrestha

202 et al., 2021). These scans are then combined to reconstruct the three-dimensional wind profiles (zonal wind u, meridional wind v, and vertical wind w) from 100 to 7000 m above ground level, 203 204 with a vertical resolution of 25-50 m and a temporal resolution of 1 s. In this study, we utilized hourly averaged horizontal wind profiles from 100 m to 2000 m at 100 m intervals above the 205 surface at Wantagh, NY, on the south shore of LI to estimate the depth of sea breeze developed 206 from the Atlantic Ocean. Details of sea breeze feature identification are described in Section 2.3. 207 The Doppler lidar site is located less than 0.5 km away from NYSM surface site WANT and 208 hereafter referred as WANT as well. 209

Weather map composites from National Oceanic and Atmospheric Administration (NOAA) High-Resolution Rapid Refresh (HRRR, https://registry.opendata.aws/noaa-hrrr-pds/) hourly reanalysis fields, with a horizontal resolution of 3 km, were used to characterize large-scale weather conditions for local circulation clusters of interest during hot summer days. In addition, we employed the gradient in HRRR hourly surface wind fields to locate the sea breeze fronts in the case studies in Section 3.3.

We also obtained Version R1 vertical column HCHO and NO₂ densities below the aircraft 216 (resolution of 0.01°) retrieved from The Geostationary Coastal and Air Pollution Events (GEO-217 CAPE) Airborne Simulator (GCAS) and Geostationary Trace gas and Aerosol Sensor 218 Optimization (GeoTASO) onboard National Aeronautics and Space Administration (NASA) 219 220 aircrafts during the LISTOS 2018 field campaign (https://www-air.larc.nasa.gov/cgi-221 bin/ArcView/listos). These observations were used to investigate the spatial distribution of HCHO and NO₂ and derive the ozone chemical regime during morning and noon times in the 222 case studies (Section 3.3) under different background circulation and sea breeze onset conditions. 223

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Figure 1. Coastal geometry of the NY metropolitan area and locations of observation stations. 227 Dash black rectangle in the main figure indicates the location of the NYC insert. Meteorological 228 sites (*Met*) from NYSM, ASOS, AQS, and NDBC are marked with black dots. Thick gray 229 outlines highlight eight NYSM sites over land and two NDBC sites over the ocean used for 230 regional quantitative analysis (Met rg) in Section 3. AQS ozone sites (Ozone) are represented by 231 red triangles. Among them, two sites in the NYC region with NOx observations (NOx) used in 232 Section 3.2 are highlighted by green reverse triangles. Locations referenced in the study are also 233 marked on the map, with "Phila" representing Philadelphia. 234

235 **2.2 Clustering of local weather conditions**

In order to study the impacts of local circulation on ozone spatial and temporal distribution, it is necessary to limit the influence of temperature given the high correlation between surface ozone level and temperature. K-means clustering is an effective and efficient technique to separate features of interest into distinctive groups (Li et al., 2020; Macqueen, 1967; Xu et al., 2021). It separates samples into clusters by minimizing the *inertia* or within-cluster sum-of-squares:

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$$\sum_{i=1}^{I} \sum_{j=1}^{J} w_{ij} \left\| x_i - \mu_j \right\|^2$$
(1)

in which x_i is the input feature vector of sample *i*, with *i* from 1 to *I*; μ_j is the centroids of cluster *j* (mean of samples in cluster *j*). *j* is from 1 to a user pre-defined cluster number *J*. $w_{ij} = 1$ when sample x_i belongs to cluster *j*; otherwise, $w_{ij} = 0$.

Major limitations of k-means clustering are that the number of clusters needs to be preselected 245 246 and the convergent results are dependent on initial centroids. We addressed these with the help of the KMeans function in scikit-learn (Pedregosa et al., 2011), a free machine learning library for 247 the Python programming language. Further details will be discussed below when clustering 248 temperature and wind conditions at QUEE. QUEE is selected because 1) as part of the newly 249 established NYSM, this site delivers quality controlled hourly average meteorological 250 observations with extraordinary temporal data coverage; 2) it is "collocated" with an active AQS 251 site reporting both ozone and NOx observations; 3) observations from AQS and NYSM at this 252 site are representative in both physical and chemical conditions around the NYC where high 253 254 ozone episodes occur frequently.

For temperature, 24 hourly mean temperatures at QUEE from 00:00 to 23:00 were directly used as input features. We preselected the number of clusters at three, and the initial cluster centroids sampling method was set to be based on an empirical probability distribution to speed up convergence and guarantee the *inertia* converging to a global minimum (Pedregosa et al., 2011). The resultant clusters are named *Hot*, *Moderate*, and *Cool* days based on their diurnal temperature profile, and their statistics will be discussed in Section 3.1.

261 We also applied k-means clustering to categorize local circulation conditions represented by six features derived from diurnal time series of hourly surface wind conditions at QUEE. The 262 263 features were adapted from Li et al. (2020), in which they clustered the wind conditions in the Houston and San Antonio regions separately based on seven characteristics derived from local 264 wind observations. The first four features of morning and afternoon wind components remain 265 similar: morning zonal (u, first feature) and meridional (v, second feature) winds were averaged 266 267 during the early morning of 04:00-06:00 and afternoon u (third feature) and v (fourth feature) winds were averaged during the early afternoon of 14:00-16:00, respectively. As there is no 268 269 apparent land breeze development during the early morning of sea breeze days in NY metropolitan area (McCabe & Freedman, 2023), days with strong southerly wind (southerly 270 271 throughout the day) and sea breeze days (calm in the early morning and southerly during sea

272 breeze hours in the region) would not be distinguishable using the original last three features in Li et al. (2020) as they all focus on the temporal variation of wind directions throughout the day 273 274 (results for QUEE shown in Figure S1). Thus, we replaced these last three features with two new ones designed to separate southerly from sea breeze days: average early morning (04:00-06:00) 275 wind speed (fifth feature) and the ratio of morning and afternoon mean wind speed (sixth feature) 276 by dividing the average early morning wind speed by the average afternoon (14:00-16:00) wind 277 speed. All features were normalized across the three summers before the cluster analysis since k-278 means clustering groups the features of interest by minimizing the Euclidean distance to the 279 centroids. Normalization allows each feature to have the same weight during the process. The 280 number of clusters was tested between 3 and 6, and four clusters yielded the best wind condition 281 separations (small *inertia* and reasonable wind clusters by visual inspection of the clustering 282 results for all summer days) were adopted. The same initialization approach was set to speed up 283 the convergence and find the global minimum as used in the temperature clustering. The four 284 resultant clusters of Sea Breeze (SB), Oscillation (O), Southerly (S), and Westerly (W) days, all 285 named based on their wind directions and speeds throughout the day, will be discussed in detail 286 287 in Section 3.1.

288 **2.3 Sea breeze feature identification**

Critical features of sea breeze impacting coastal ozone levels, as identified in previous case 289 studies, are estimated based on NYSM surface and wind lidar observations at the coastal site of 290 WANT during Hot SB days. Figure 2 demonstrates the identification of those sea breeze features 291 on Aug. 5, 2018, including sea breeze depth (SB_Depth, Figure 2a), sea breeze onset v wind 292 instant change (SB_dV, Figure 2b), sea breeze onset time (SB_OnT, Figure 2b), and sea breeze 293 strength (SB_Str, Figure 2b). SB_dV is defined as the maximum hourly gradient (change in the 294 next hour from the current hour) in 10 m v wind during 07:00-16:00. SB OnT (10:00 in the 295 example) is the hour following the time when SB_dV occurs. SB_str is estimated to be the 296 297 average 10 m v difference between 3 hours after onset time (11:00-13:00 in the example) and early morning of 04:00-06:00. SB_Depth is the height above ground at which the correlation of v 298 299 wind with 10 m v wind drops below 0.5 starting from 100 m to 2000 m. Six hours (starting from the hour when SB_dV occurs, 09:00-14:00 in the example), instead of the whole sea breeze 300 301 period, were utilized in the estimation to avoid the ambiguity in deciding the sea breeze cessation

time and the impact of low-level jet developed in the later hours during some of the sea breeze days. Note that the detected SB_OnT and SB_Str might not be as reliable when the wind speed of the Atlantic sea breeze increases gradually on some of the days with relatively stronger background onshore flow in the early morning. In addition, stronger upper-level flow and lidar data availability during some of the *Hot SB* days might challenge the quality of the identified SB_Depth .

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Figure 2. Sea breeze feature identification at WANT on Aug. 5, 2018: (a) 10 m surface wind and horizontal wind profiles at 100s meters above ground level with identified sea breeze depth (SB_Depth) marked with a thick black line; (b) time series of 10 m v wind (negative values indicate wind from the north and positive values indicate wind from the south; the same convention applies throughout this study) and its gradient with sea breeze strength (*SB_Str*), onset time (*SB_OnT*), and onset v wind instant change (*SB_dV*) illustrated.

Sea breeze front locations in two case studies in Section 3.3 were estimated based on hourly 316 HRRR surface wind reanalysis fields by calculating the wind gradient Grad in x and y directions 317 as follows: 318

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$$Grad[x, y] = \sqrt{(u[x, y] - u[x + 1, y])^2 + (v[x, y] - v[x + 1, y])^2} + \sqrt{(u[x, y] - u[x, y + 1])^2 + (v[x, y] - v[x, y + 1])^2}$$
(2)
which x and y represent the grid cell location; u and v are the 10 m wind components. The

322 in w le 323 longitude *lonxy* and latitude *latxy* of the grid cell for *Grad* is calculated by:

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$$lonxy[x, y] = \frac{lon[x,y] + lon[x,y+1]}{2}$$
 (3)

325
$$latxy[x, y] = \frac{lat[x, y] + lat[x+1, y]}{2}$$
 (4)

in which lon and lat are the longitude and latitude of the original HRRR grid cell. Higher values 326 327 of Grad indicate higher contrast in wind speed and directions. Lines formed by connected higher Grad values would signal the location of possible fronts, including sea breeze fronts. 328

3 Results and Discussions 329

3.1 Meteorological clustering results and summertime all-weather ozone distribution 330

The clustered diurnal cycles of temperature at QUEE are shown in Figure 3. During the three 331 summers (median temperatures are described below), 70 days are clustered as Hot days with 332 maximum temperature around 30.7°C at 14:00 and minimum temperature around 24.2°C at 333 05:00, 132 days are clustered as Moderate days with maximum temperature around 26.8°C at 334 14:00 and minimum temperature around 21.3°C at 05:00, and 72 days are clustered as Cool days 335 336 with maximum temperature around 22.2°C at 15:00 and minimum temperature around 17.2°C at 05:00. The clean separation among the interquartile ranges of the three temperature clusters 337 indicates the effectiveness of the k-means clustering. 338

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Figure 3. Diurnal cycles of three temperature clusters of *Hot*, *Moderate*, and *Cool* days at QUEE.
Median temperatures (solid lines) with their maximum and minimum (marked in dots) and
interquartile range (light-colored shades) of each cluster are shown. The number of days in each
cluster is shown in the legend.

Figures 4a and b display the average hourly wind vectors and early morning and afternoon wind 346 roses for four wind clusters of Sea Breeze (SB), Oscillation (O), Southerly (S), and Westerly (W) 347 at QUEE. In total, 113 SB days have been identified with light wind in the early morning and 348 much stronger southerly wind from the Atlantic Ocean in the afternoon. There is no apparent 349 land breeze developed during nighttime, agreeing to what is found in McCabe & Freedman 350 (2023). The weaker sound breeze developed from LIS towards LI and CT, as identified in 351 previous studies (Han et al., 2022; Novak & Colle, 2006), normally does not penetrate into NYC 352 and are thus not reflected in the wind conditions at QUEE. Similar to the clustering results in Li 353 et al., 2020, an Oscillation cluster has been identified. It consists of 69 days impacted by strong 354 large-scale wind flow with slowly clockwise-rotating wind. The Southerly cluster represents 36 355 days with consistent southerly wind throughout the day. Its daytime southerly wind tends to be 356 strengthened by sea breeze due to the land-sea temperature contrast, and its wind speed peaks 357 early around noontime compared with 15:00 in the SB cluster with light synoptic flow in the 358 region. Note that sea breeze, any wind that blows from a large body of water toward or onto a 359 landmass, as a physical phenomenon might happen under many synoptic-scale wind conditions, 360 361 such as in the Southerly cluster; the italicized cluster Sea Breeze (SB) only refers to the days in 362 that cluster with relatively calm background wind fields. Lastly, 57 Westerly days have been found to have constant westerly wind blowing into NYC. The westerlies are intensified by a 363 364 northwest component during the afternoon.



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Figure 4. Wind clusters at QUEE. (a) diurnal cycles of mean wind for four wind clusters of *Sea Breeze (SB), Oscillation (O), Southerly (S),* and *Westerly (W)*; (b) early morning (04:00-06:00,
top panels) and afternoon (14:00-16:00, bottom panels) wind roses for each wind cluster. The
number of days in each cluster is shown in (b).

With temperature and wind diurnal cycles clustered, we first confirmed the effectiveness of those 372 373 clustering by checking the consistency of their statistics across ten combined clusters (Table 1, two clusters with less than six days are excluded from the analysis due to the scarcity of such 374 375 weather conditions). As expected, daytime temperature is generally consistent during *Hot* (28.0-30.0°C), Moderate (24.9-25.7°C), and Cool (19.9-21.6°C) days across different wind conditions; 376 morning wind speed and wind speed ratio statistics are also consistent at different temperature 377 levels. For example, SB days have the lowest morning wind speeds at 1.7-1.8 m/s and ratios of 378 379 morning and afternoon wind speeds of 0.4 for all three temperature clusters.

Summertime all-weather ozone distribution for all combined meteorological clusters is shown in 380 Figure 5 and summarized in Table 1. It is clear that temperature and local circulation have 381 distinctive impacts on daily ozone levels in the region. As noted previously, higher temperature 382 facilitates higher ozone production in general. A specific wind pattern exhibits consistent effects 383 on the spatial distribution of ozone, regardless of temperature levels (also see Figure S2 for 384 DMA8 spatial anomaly with the cluster average removed). Specifically, SB days favor high 385 386 ozone levels along the Philadelphia-NYC-LIS corridor (dashed gray line in Figure 5a), especially for Hot and Moderate days (Figures 5a and d, S2a and d). During S days, the region is occupied 387

by moist and modest marine air brought in by the steady southerly flow. Due to the similarity in wind directions compared with *SB* days, similar high ozone locations are observed on *S* days, with an expansion of high ozone levels into the CT inland region (Figures 5b and f, S2b and f). The region south of the Philadelphia-NYC-LIS corridor experiences higher ozone levels during *W* days at all temperature levels (Figures 5c, g, and j; S2c, g, and j). *O* days tend to be accompanied by lower temperature levels (*Moderate* and *Cool*) and observe the lowest ozone levels during *Cool* summer days with an average DMA8 of 33.4 ppb (Figures 5e and i; Table 1).

It is clear that the most polluted days are closely associated with *Hot SB* days (DMA8 at 60.4 ppb and 8.7 sites with ozone exceedances per day on average) with weak large-scale flow in the early morning, followed by *Hot W* days (DMA8 at 53.5 ppb and 2.4 sites with ozone exceedances per day). *Hot* summer days during which ozone exceedances (DMA8 > 70 ppb) are expected to happen much more frequently will be further studied in the following sections.

Table 1. Summary of summertime meteorological clusters and regional ozone characteristics. SD stands for Standard Deviation. Daytime temperature is averaged from 07:00 to 17:00 EST. Morning wind speed and wind speed ratio refer to the fifth and sixth feature, respectively, as discussed in Section 2.2. Ozone exceedance (DMA8 > 70 ppb) occurrences count all sites, as shown in Figure 5, during all days in each cluster. Ozone exceedance occurrences (/day) are then calculated by dividing that by the number of days in the cluster.

	Temperature cluster	Hot			Moderate				Cool		
	Wind cluster	SB	S	W	SB	0	S	W	SB	0	W
	Number of days	37	7	22	54	27	25	26	22	37	9
Cluster	Daytime temperature mean (°C)	29.4	28.0	30.0	25.7	24.9	25.4	25.5	21.1	19.9	21.6
statistics at	Daytime temperature SD (°C)	1.2	1.3	1.9	1.3	1.6	1.1	1.4	1.7	2.5	1.5
QUEE	Morning wind speed mean (m/s)	1.8	3.0	2.6	1.7	2.5	3.9	3.3	1.7	3.3	3.3
	Morning wind speed SD (m/s)	0.4	0.5	0.8	0.5	0.9	1.3	1.1	0.6	1.0	1.0
	Wind speed ratio mean	0.4	0.7	0.6	0.4	0.9	0.8	0.7	0.4	1.1	0.6
	Wind speed ratio SD	0.1	0.4	0.2	0.1	0.3	0.3	0.2	0.1	0.4	0.2
	DMA8 spatial mean (ppb)	60.4	46.9	53.5	49.1	41.2	42.9	41.7	41.2	33.4	40.8
Pagional	DMA8 spatial SD (ppb)	5.5	5.0	6.0	3.1	4.0	2.8	3.8	2.3	1.6	2.9
ozone	Ozone exceedance occurrences	323	9	52	69	2	32	7	0	0	0
	Ozone exceedance occurrences (/day)	8.7	1.3	2.4	1.3	0.1	1.3	0.3	0.0	0.0	0.0

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Figure 5. Average DMA8 for all meteorological clusters with at least six days during the
summertime (JJA) of 2017-2019. The number of days in each cluster is marked on the map.

412 **3.2 Impacts of local circulation on ozone during hot summer days**

As sea breeze develops and penetrates NYC during both Hot SB and S days, we first examined 413 whether sea breeze developed during *Hot W* days along south LI by looking at wind conditions at 414 WANT. It turns out sea breeze develops along south LI during most Hot W days but fails to 415 416 penetrate into NYC (QUEE) because of the dominant westerlies. Thus, Hot W days are further divided into two subsets according to the afternoon (13:00-16:00) mean v wind direction at 417 418 WANT (Figure 6). Namely, they are Hot W_w and Hot W_sb days, with lowercase w and sb representing afternoon conditions at WANT. Five W_w days with persistent strong 419 420 northwesterlies at WANT in the afternoon are detected among 22 Hot W days and will be ignored hereafter because of the limited number of days and their low ozone levels. Cluster W_sb 421 consists of 17 Hot W days with delayed sea breeze onset (around 11:00 at the coastal site of 422 WANT, as in Figure 6, compared with 09:00-10:00 at QUEE, a site further inland in NYC, as in 423

Figure 4a) and penetration of sea breeze inland to the northwest part of the study region (such asOUEE) intercepted by the dominant westerlies.



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Figure 6. Diurnal cycles of mean 10 m wind at WANT during *Hot W* wind subsets of W_w (5 days) and W_{sb} (17 days). The lowercase *w* and *sb* in the subset names represent coastal wind conditions of westerly and sea breeze at WANT.

Weather maps based on composites of HRRR reanalysis fields of 850 hPa geopotential height 430 and wind during Hot day clusters (SB, W_sb, and S) are shown in Figure 7. It should be noted 431 that sea breeze developed from the Atlantic Ocean in all three clusters. However, the 432 characteristics of the sea breeze vary in geographical extent, magnitude, and timing. The 433 Bermuda High pressure system ("H" in Figure 7) sits near the Southeast Coast of the 434 435 contiguous United States (CONUS) during summertime, generating light southwesterlies in the study region during Hot SB days (Figures 4a and 7a). When the low-pressure system in 436 southeastern Canada ("L" in Figure 7) moves eastwards and pushes the north part of Bermuda 437 High south during Hot W sb days, westerly components of the local wind around the NY 438 439 metropolitan area strengthen, delaying the sea breeze development near the coast and preventing its penetration further inland to the city center (Figures 4a, 6, and 7b). Strong southerly wind in 440 the New York Bight and its coastal region is generated during *Hot S* days as the Bermuda High 441 advances closer to the Northeast Coast of CONUS with its large-scale clockwise rotation (Figure 442 443 7c). The southerly wind strengthens during daytime as the Atlantic sea breeze develops (Figure

444 4a).



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Figure 7. 850 hPa average weather conditions at 07:00 for *Hot* (a) *SB* (37 days), (b) W_sb (17 days), and (c) *S* (7 days) clusters. Wind speed and direction are represented by black vectors; colored contours indicate the 850 hPa geopotential heights in geopotential meters (gpm) with intervals of 10 gpm; relative locations of major high (H) and low (L) pressure systems impacting the NY metropolitan area are marked in white.

Similar to what is shown in Figure 5, Figures 8a-c displays the spatial distribution of average 452 DMA8 for three Hot day clusters, with the modification of W to W_sb days. As expected, mean 453 DMA8 on W sb days (median value of 56.0 ppb) increased notably along LI and coastal CT 454 compared with *Hot W* days (median value of 53.8 ppb) in Figure 5 after removing the *W_w* days. 455 In addition, the probability of daily DMA8 exceeding the NAAQS level of 70 ppb (Pex) for each 456 cluster is shown in Figures 8d-f. During Hot SB days, the three most polluted sites located in 457 coastal Fairfield County (southwest CT) observed mean DMA8 higher than 70 ppb with 458 459 corresponding Pex around 50%. Among them, mean DMA8 and Pex at Westport, CT, is 71.5 ppb and 43.2%, respectively. Slightly lower DMA8 and *Pex* are observed at other sites along the 460 CT coastline and Queens, NY (mean DMA8 = 63.8 ppb, Pex = 40.5%). Pex for W_sb (median 461 value of 5.9%) and S days (median value of 0.0%) and DMA8 for S days decreased dramatically 462 463 compared with SB days (median value of 18.9%).



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Figure 8. (a-c) DMA8 mean and (d-f) percentage of days with DMA8 exceeding NAAQS of 70 ppb (*Pex*, sites with no exceedance are not shown) on *Hot* (a, d) *SB*, (b, e) W_sb , and (c, f) *S* days. Note that sites with no ozone exceedance (*Pex* = 0.0) are not shown in (d-f) for clarity, but the minimum, median, and maximum statistics are calculated based on all sites.

To better understand the change in DMA8 levels for different local circulation clusters, we 469 looked into the regional ozone, NOx (NYC only where observations are relatively abundant), 470 and temperature diurnal cycles (Figure 9), with differences in DMA8 between most polluted Hot 471 SB and W sb days shown in Figure 9h. In addition, ozone tendency maps overlaid with local 472 wind conditions are displayed in Figure 10. The three high ozone exceedance probability (high 473 *Pex*) regions of NYC (squares in Figure 9h), CTcoastal (coastal region of western CT, triangles 474 in Figure 9h), and LICTe (LI and coastal region of eastern CT, reverse triangles in Figure 9h) are 475 selected and grouped manually based on their spatial proximity and similarity in the 476 characteristics of their ozone diurnal cycles during SB, W_sb, and S days. 477

478 When sea breeze development in the New York Bight is delayed and its penetration into the 479 NYC and shores of western LIS is intercepted by the dominant westerlies (W_{sb}) (Figures 10d-f), DMA8 in the hot spots around NYC and western CT would typically drop 9-10 ppb (Figure 9h) 480 under comparable temperature levels (Figures 9b, e, and g, Table 2) compared with *Hot SB* days. 481 482 The average regional decrease of DMA8 for NYC and CTcoastal regions is 6.7 and 8.3 ppb, 483 respectively (Table 2). Higher levels of local NOx emission are accumulated in the NYC region during the morning time of SB days due to the continuous NOx emission from industrial sources 484 and stagnant nighttime conditions. NOx levels begin to drop after 07:00 as the photochemical 485

486 production of ozone is initiated after sunrise (Figures 9a and c). Compared with SB days, the rate of ozone increasing in the early morning during W sb days is much lower, and ozone level 487 488 plateaus much earlier, around 11:00, at lower levels in NYC as strong westerly carry the local pollutants downwind (Figures 9a and 10d-e). In CTcoastal, ozone increases fast in the morning 489 time and decreases in the afternoon at a similar rate during both SB and W sb days (Figure 9d); 490 the 9-10 ppb difference in DMA8 in CTcoastal stems from the additional fueling during SB days 491 around noon time. While ozone build-up time during W_{sb} days in LICTe becomes hours longer 492 compared with SB days because of the stationary sea breeze/westerly front in the region (Figures 493 10d-f), comparable or even higher DMA8 levels are observed, especially in southeastern LI 494 (Figure 9h). Zhang et al. (2022) also found that the high ozone cases at the Heckscher State Park 495 (south shore of LI) always involve the westerlies and southwesterly flow. 496

497 Lowest ozone levels are observed during Southerly days for all three regions because of the presence of marine air (warmer during nighttime and colder during daytime) throughout the day 498 499 brought in by the steady southerly flow (Table 2; Figures 9 and 10g-i). The largest difference in 500 ozone levels is observed in LICTe due to its proximity to the northerly moving and relatively 501 cleaner marine air from the Atlantic Ocean. Ozone starts to drop as early as 12:00 at some sites in NYC and LICTe, especially for the highly polluted site at Queens, NY (Figures 9a, f, and 502 Figure 10h). CTcoastal experiences the least improvement in air quality during S days, likely due 503 to the transported polluted air trapped in the shallow boundary layer above LIS (Torres-Vazquez 504 505 et al., 2022). The secondary peak of ozone in the late afternoon, around 19:00-20:00, during Hot 506 S days might be brought to all three regions by the return flow of the sea breeze circulation, with 507 southerly flow dying down in the meantime.

Table 2. *Hot* summer day average condition for NYC, CTcoastal, and LICTe. Daytime
 temperature is averaged from 07:00 to 17:00 EST.

		NYC			CTcoastal			LICTe	
	SB	W_sb	S	SB	W_sb	S	SB	W_sb	S
Daytime Temperature (°C)	30.0	30.2	28.6	28.3	28.8	27.2	27.8	28.4	27.2
DMA8 (ppb)	63.7	57.0	49.1	67.6	59.3	51.7	62.5	65.6	45.3

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Figure 9. Regional ozone, surface wind, temperature, and NO_x (NYC only) diurnal cycles during *Hot SB*, W_sb , and *S* days for (a-c) NYC (square in h), (d-e) CTcoastal (triangle in h), and (f-g) LICTe (reverse triangle in h), with stations used for the regional composite marked in (h) the *Hot SB* and W_sb DMA8 difference map with number of sites in each region shown in the legend. Solid lines represent average conditions (median condition for ozone) with interquartile ranges shaded in the same colors.

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Figure 10. Average ozone tendency, calculated by subtracting ozone at the current hour from ozone at the next hour, during *Hot* (a-c) *SB*, (d-f) W_sb , and (g-i) *S* days from 11:00 to 13:00 (left to right). Sea breeze front locations are manually marked with dashed gray lines (the weaker sound breeze towards LI is not marked with limited surface observations) based on the observed surface wind (magenta wind vectors).

529 **3.3** How local circulation modulate ozone exceedances in extreme heat

With the projected extreme heat in NYC increasing in intensity and duration throughout the 21st 530 531 century (RCP4.5 and RCP8.5) (Ortiz et al., 2019), ozone levels are expected to rise under current emission conditions. This section examines whether local circulation can modulate ozone 532 533 exceedances in extreme high heat by analyzing the meteorological characteristics for ozone exceedance and non-exceedance days at the most polluted NYC (Queens, NY as example) and 534 535 CTcoastal (Westport, CT as example) regions (Figures 8a and d) during hot SB days. In addition to four sea breeze features identified at WANT, as discussed in Section 2.3, we also included the 536 537 following meteorological features:



- monitoring site (NYSM site QUEE for Queens, NY, and co-located wind observation at
 AQS site 090019003 for Westport, CT).
- Regional average early morning v wind during 04:00-06:00 (V_am) at eight NYSM surface sites highlighted with thick black outlines in Figure 1.
- Regional average early morning u wind during 04:00-06:00 (U_am) at the same eight sites.
- Regional average daytime mean temperature during 07:00-17:00 (T) at the same eight sites.
- Regional average daytime temperature difference from the sea during 07:00-17:00 (*dT*).
 Sea temperature is calculated at two NDBC sites highlighted with thick blue outlines in
 Figure 1.

551 To select those days with extreme high heat and minimize the ozone dependence on temperature, 552 we limited our analysis to 16 Hot SB days with regional daytime temperature (T) higher than 29°C, defined as extreme heat/hot days in this study. The threshold of 29°C is manually picked 553 to effectively de-correlate DMA8 at Queens, NY, and T based on their scatter plot (Figure S3). 554 555 Note that ozone observations at Westport, CT are missing during 2 out of the 16 days. As 556 expected, ozone exceedances (DMA8 > 70 ppb) happened during the majority of the extreme hot days at both locations (11 out of 14 days for Westport, CT, and 11 out of 16 days for Queens, 557 NY). Among the 14 days with ozone observations at Westport, CT, 7 days witnessed ozone 558 exceedances at both locations. Only 1 (Aug. 5, 2018, Sunday) out of the 14 days did not 559 560 experience ozone exceedance for both locations, most likely due to the limited availability of NOx emissions during the weekend (Figure S4). DMA8 that day for Westport and Queens are at 561 562 60.6 ppb and 66.4 ppb, respectively. Days with different ozone exceedances at the two locations are Jun. 30, 2018, and Jun. 28, 2019, for ozone exceedances at Queens but not Westport; and Jul. 563 19, 2017, Aug. 28, 2018, Jul. 28, 2019, and Jul. 30, 2019, for ozone exceedances at Westport but 564 565 not Queens.

Figure 11 presents the ozone levels and meteorological features (detailed in Section 2.3 and the beginning of this section) during ozone exceedance and non-exceedance days at Queens, NY, and Westport, CT. Panels from left to right are organized by units and magnitudes of the variables. Even though Queens is located closer to major emission sources, its DMA8 and hourly 570 peak ozone level (O_3_max) are significantly lower than those of Westport during ozone 571 exceedance days. In addition, both DMA8 and O_3 max at Westport differ significantly during 572 ozone exceedance and non-exceedance days. Both phenomena signify the importance of transport in modulating ozone exceedances in the region (Figures 11a and g). During extreme hot 573 days, the median T difference between ozone exceedance and non-exceedance days is very 574 minimal at 0.1 and 0.2°C at Queens and Westport, respectively (Figures 11e and k). The same is 575 576 true for the depth of sea breeze (SB_Depth) (Figures 11d and j). Likely due to the removal of colder days, no significant difference is found in median sea breeze onset (SB OnT) and arrival 577 time (SB_ArrT) (Figures 11b and h). Sea breeze tends to develop early in coastal areas during 578 extreme hot days because of the higher contrast in land/sea temperature and penetrated into NYC 579 580 two to three hours later. Due to the impact of the LIS sound breeze, there is large uncertainty in 581 the estimation of sea breeze arrival time *SB_ArrT* at Westport, CT.

582 Theoretically, a higher contrast in land/sea temperature (dT) observed during ozone exceedance 583 days (around 0.4°C higher than non-exceedance days for both locations, Figures 11f and 1) would favor a stronger sea breeze onset. However, the sea breeze strength (SB_Str) and onset 584 585 type (higher SB_dV for abrupt sea breeze onset and lower for gradual sea breeze onset) depend 586 more on the early morning wind conditions in the region (Figures 11c and i). During extreme hot SB days with relatively stagnant early morning conditions, a relatively stronger southerly wind 587 component (higher V_am around 1m/s) in the early morning will favor gradual (lower SB_dV) 588 589 weak (lower SB Str) sea breeze development (Figure 11c). When coupled with slightly stronger 590 morning westerlies (higher U_am), these gradual weak sea breeze days would typically yield 591 lower chances of ozone exceedances in NYC (Figures 11a and c). On the contrary, a relatively stronger northerly wind component (higher negative V am around -1 m/s) in the early morning 592 will likely lead to abrupt (higher SB_dV) strong (higher SB_Str) sea breeze onset (Figure 11i). 593 594 This finding is consistent with the previous lake breeze study in Salt Lake Valley, in which they concluded that opposing flow would contribute to stronger lake breeze front (Blaylock et al., 595 2017). In this case, the south shore of CT would see a lower possibility of ozone exceedances, 596 597 especially when morning westerly flow is weaker (lower U am) (Figures 11g and i). In the following, we will illustrate how these differences in local circulation conditions, stemming from 598 the subtle early morning meridional wind (V_{am}) contrast, modulate ozone exceedances in these 599 two regions using two case studies during the LISTOS field campaign. 600

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Figure 11. Ozone levels and meteorological features during ozone exceedance days (DMA8 > 603 70 ppb, red) and non-exceedance days (DMA8 \leq 70 ppb, blue) for Queens, NY (top panels) and 604 Westport, CT (bottom panels) under extreme high heat (T > 29 °C). Panels from left to right are 605 for: (a, g) DMA8 and O_3 max with NAAOS level of 70 ppb marked with dash line; (b, h) 606 SB_OnT and SB_ArrT with 12:00 marked with dash line; (c, i) wind conditions of SB_dV, SB_Str, 607 U am, and V am with 0 m/s marked with dash line; (d, j) SB Depth; (e, k) T; and (f, l) dT. 608 Boxplots show interquartile range (IQR) with the median value in black; the whiskers spread to 609 the farthest point if within 1.5 times the IQR from the box; data points outside the whisker limits 610 are shown as black open circles. 611

612 Case 1: Abrupt strong sea breeze onset

Jun. 30, 2018, is on a Saturday with regional daytime mean temperature (T) of 30.2° C and 613 northerly background wind in the early morning (V am = -1.0 m/s). Higher ozone levels around 614 noon time are confined within the NYC area (Figure 12b). Specifically, DMA8 at Queens (ozone 615 exceedance) and Westport (non-exceedance) are at 72.6 and 60.0 ppb, respectively. The case 616 study is demonstrated using monitored surface wind and temperature (Figure 12a), ozone and 617 NO₂ (Figure 12d), and spatial distribution of monitored ozone (Figures 12b and c) and ozone 618 tendency (Figures 12e and f) overlaid with surface wind gradient based on HRRR. In addition, 619 620 the vertical column density (VCD) of NO₂ and HCHO, as well as the chemical regime estimated by their ratio, are shown in Figure 13 to illustrate chemical conditions during the early morning 621 622 and afternoon.

623 In the early morning, around 05:00, the ozone level at Queens dipped below 5 ppb, indicating strong NO titration (O₃ removal by NO) around NYC because of continuous NO emission from 624 625 industrial sources and stagnant conditions during the nighttime, which facilitated the NOx accumulation around the source region (Figures 12a and d). This type of extensive urban-scale 626 627 "hole" with extremely low surface ozone has been widely reported in previous studies such as Zhang et al. (2004). After sunrise and before the sea breeze arrival, the early morning 628 629 northwesterly wind strengthened in NYC (Figure 12d), transporting NO₂ in the southeastern direction (Figure 13a). LIS sound breeze towards coastal CT developed at 10:00 as the land 630 surface temperature rose. It remained a southerly flow throughout the day, making ozone over 631 coastal CT sensitive to ozone and its precursors over the LIS. The sea breeze fully developed at 632 WANT around 10:00 and arrived at QUEE only one hour later (Figure 12a). Apart from the sea 633 breeze from the Atlantic Ocean, complex local circulation systems around the NYC region 634 developed and sustained into the afternoon due to the urban/suburban and land/water temperature 635 636 contrast (Figures 12e and f). These systems include the divergent sound breeze along LIS, river breeze along the Hudson River, and bay breeze at the coast of Upper Bay, Lower Bay, and 637 638 Newark Bay, and convergence tendency (UHI) around urban centers of NYC and Newark. After the sea breeze penetrated into the NYC area, its further north movement was hindered by these 639 640 strong local circulation systems. NO₂ transported southeast of NYC was recircuited back with the sea breeze from the Atlantic Ocean and accumulated around the city center (Figure 13d) 641 642 because of the stationary convergent flow generated from the systems described above, facilitating long hours of ozone buildup around the city center, especially north of NYC at White 643 Plains (Figures 12f). Figure 12d highlights the long hours of high ozone observed at Queens. 644

Although the high ozone levels tend to be confined around NYC in abrupt sea breeze cases, the 645 complex circulation systems do not always fully develop because of the difference in synoptic 646 wind conditions, in addition to urban/suburban and land/water temperature contrast (Figure S4 647 and S5). However, sound breeze towards CT and LI is consistently detected in all three cases 648 649 (Figures 12, S4, and S5). Without the complex circulation system as on Jun. 30, 2018, sea breeze typically arrives at Queens two hours after its onset at WANT with light northwesterlies in the 650 651 early morning. Ozone at Queens, sitting in between two fronts moving towards each other (front 652 parallel to the CT coastline exserted by the background northwesterlies and sea breeze front from

the Atlantic Ocean), rises faster before sea breeze from the Atlantic Ocean arrives and dropsimmediately after the sea breeze front passes by in two other cases shown in Figure S4 and S5.

Ozone levels in CTcoastal, however, are more NOx-sensitive. In days with light northwesterly flow in the early morning, NOx from NYC is not effectively transported to LIS and coastal CT, making the regional ozone production fall in the NOx-limited regime throughout the day (Figures 13c and f). Balanced between transport and chemical reactions, ozone at Westport plateaued early from 11:00 likely due to the limited NOx (Figures 12 and 13c, f), till night time at around 20:00 (similar in two other cases in Figure S4 and S5).

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Figure 12. Abrupt strong sea breeze onset case on Jun. 30, 2018: (a) 10 m v wind at WANT, 663 664 OUEE, and Westport with diurnal mean 2m temperature in the NYC-LI region; (b-c) Ozone 665 distribution (filled color circles) and surface wind vectors from observations (magenta) and 666 HRRR reanalysis wind field (black vectors) at 12:00 (c) and 14:00 (d) with front locations (green shade) signified by surface wind gradient (Grad, see Section 2.3); (d) Ozone and NO₂ time series 667 at Queens and Westport with NAAQS level of 70 ppb marked with dash line; (e-f) hourly ozone 668 tendency (filled color circles) and local wind conditions (same as in b) around NYC at 12:00 (e) 669 and 14:00 (f). 670

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Figure 13. Abrupt strong sea breeze onset case on Jun. 30, 2018: vertical column density (VCD) of (a,d) NO₂, (b,e) HCHO, and (c,f) chemical regime estimated by their ratio (HCHO/NO₂) below aircraft (data from GeoTASO) in the morning (a-c) and early afternoon (d-f) around NYC. The ozone chemical regime in (c) and (f) follows the method used in Tao et al. (2022) with thresholds estimated in Jin et al. (2020) (transitional regime when HCHO/NO₂ ratios range from 2.9 to 3.8, with ratios below this range for the NOx-saturated/VOC-limited scenario and above for the NOx-limited/VOC-saturated scenario).

681 Case 2: Gradual weak sea breeze onset

Aug. 28, 2018, is on a Tuesday with regional daytime mean temperature (T) of 31.5° C and 682 southerly background wind in the early morning ($V_{am} = 0.9$ m/s). High ozone levels around 683 noon time are observed along the CT coastline (Figure 14c). Specifically, DMA8 at Queens 684 (ozone non-exceedance) and Westport (exceedance) are at 66.3 and 84.5 ppb, respectively. 685 Noticeable higher levels of HCHO are observed in the morning and afternoon across the region 686 (Figures 15b and e), likely due to the acceleration of biogenic VOC emissions driven by the 687 higher temperature throughout the day. The higher temperature would also yield more active 688 photochemistry. The higher energy demand and weekday/weekend effect (the Jun. 30 case is on 689 Saturday, and the Aug. 28 case is on Tuesday) might also increase the NOx loading on this day. 690

691 With southwesterly background flow starting before midnight, a large load of NO_2 emitted from 692 sources around NYC was diluted downwind along LIS, making these regions NOx-saturated in 693 the morning (Figures 15a and c). NO titration is observed at Westport with very minimal levels 694 of ozone and high levels of NO_2 before sunrise (Figure 14c). Gradual (Figure 14a) and weak 695 (lower wind gradient as seen in Figures 15b, c, e, and f) sea breeze front from the Atlantic Ocean developed early and carried higher concentrations of NOx and HCHO northward with it, 696 697 favoring the extreme high ozone production in coastal CT, LI, and LIS around early afternoon (Figures 14f and 15d, e). In the meantime, the LIS sound breeze towards coastal CT developed 698 and remained a southerly flow throughout the day as in the abrupt strong sea breeze case. The 699 resultant peak hour ozone at Westport is well above 100 ppb. Likely coupled with the incoming 700 polluted air confined in the shallow marine boundary layer above LIS with very active ozone 701 production, the hourly increase of ozone at Westport could reach up to more than 40 ppb 702 (reported on Jul. 30, 2019, Figure S6) as marine air progressed north from LIS (Figure S6d). 703 Finally, the sea breeze front arrived at Westport in the late afternoon (around 17:00), and the 704 705 decreasing of ozone slowed down with its progression towards coastal CT (Figures 14, S6, and S7). 706

Likely due to the limited availability of NOx, which is diluted away by the consistent stronger southwesterly flow (Figures 12a and 14a) and shorter period of weaker convergent flow around NYC (Figures 12e and 14e) compared with abrupt strong sea breeze days, ozone at Queens was not able to achieve high peaks and quickly decreased as the sea breeze front passed by (Figure 14d). With a stronger westerly component in the background wind, the ozone level might start to drop even before the sea breeze front arrives (Figure S7).



Figure 14. Same as Figure 12, but for gradual weak sea breeze onset case on Aug. 28, 2018.



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Figure 15. Same as Figure 13, but for gradual weak sea breeze onset case on Aug. 28, 2018 (data from GCAS).

720 **4 Conclusions**

This study comprehensively investigated the impacts of local circulation on coastal ozone 721 pollution in the NY metropolitan area using three summers (2017-2019) of meteorological and 722 chemical observations. We adopted the k-means clustering technique to cluster NYC diurnal 723 temperature profiles and wind conditions. The meteorological conditions were grouped into three 724 temperature clusters of hot, moderate, and cool days and four wind clusters of sea breeze, 725 oscillation, southerly, and westerly days. It is found that specific wind patterns exhibit consistent 726 727 effects on the spatial distribution of ozone regardless of temperature levels. For example, sea breeze days favor high ozone levels along the Philadelphia-NYC-LIS corridor at both hot and 728 729 moderate temperatures. The most polluted days are closely associated with classic hot sea breeze days with weak large-scale flow with a regional mean DMA8 of 60.4 ppb and 8.7 sites with 730 731 ozone exceedances per day on average, followed by hot westerly days with DMA8 of 53.5 ppb and 2.4 sites with ozone exceedances per day. 732

The spatial and temporal features in ozone time series with respect to local circulation conditionsduring hot summer days and how local circulations contribute to such variations are further

735 studied. When sea breeze development in the New York Bight is delayed and its penetration into the NYC and northwestern LIS is intercepted by the dominant westerlies, DMA8 in some ozone 736 737 hot spots of NYC and the south shore of CT would drop 9-10 ppb under comparable temperature levels. The average regional decrease of DMA8 for NYC and coastal CT is 6.7 and 8.3 ppb, 738 respectively. Meanwhile, ozone build-up time in LI and eastern CT becomes hours longer, and 739 DMA8 increases by 3.1 ppb on average because of the stationary sea breeze/westerly front along 740 the coast. The overall probability of DMA8 exceeding the NAAQS level of 70 ppb is also 741 dramatically decreased from 18.9% to 5.9%. Ozone exceedance is highly unlikely with marine 742 air present in the region during hot southerly days. 743

744 Furthermore, meteorological characteristics most critical in modulating ozone exceedances in NYC and the south shore of CT during extreme hot sea breeze days are identified and analyzed 745 746 with case studies during the LISTOS field campaign. As expected, ozone exceedances are very common. During these days, a relatively stronger northerly wind component (offshore) in the 747 748 early morning will likely lead to an abrupt strong sea breeze onset and contribute to a lower 749 possibility of ozone exceedances on the south shore of CT. On the contrary, a relatively stronger 750 southerly wind component (onshore) in the early morning will favor gradual weak sea breeze 751 development and would typically yield lower chances of ozone exceedances in NYC. Ozone level in the NYC region (Queens, NY) increases before the major sea breeze front from the 752 Atlantic Ocean arrives in both scenarios and typically drops immediately after the sea breeze 753 754 front passes by. However, ozone level continues to rise and remains at high concentrations when 755 the further north movement of the sea breeze is hindered by a long period of convergent flow 756 into the NYC generated by complex local circulation systems such as urban heat island. On the 757 south coast of CT (Westport, CT), ozone is more NOx-sensitive. Ozone plateaus early from around 10:00-11:00 till late afternoon around 18:00-19:00 due to the limited NOx transported 758 759 into the region and lower local emission levels in the abrupt strong sea breeze case. When abundant NOx is brought downwind into the LIS and coastal CT in the gradual weak sea breeze 760 case, high spikes of ozone around noontime are observed, leading to much higher ozone peaks 761 762 and DMA8 levels in the region.

In summary, this three-summer study comprehensively examines the impacts of local circulationon the coastal ozone problem in the NY metropolitan area. It concludes how surface ozone is

impacted spatially and temporally under different temperature scenarios: all summer days, hot summer days, and extreme heat. It will serve as a reference for future case studies and provide helpful guidance for future monitor deployments or field campaigns in the region. However, due to the limited vertical observations in such a long timeframe, this study does not cover ozone vertical profiles.

Despite being extensively studied, coastal urban environments are complicated and warrant 770 further investigations with high-resolution monitoring and modeling. In this study, only one site 771 for each of the most polluted regions of NYC and the south coast of CT is studied due to the lack 772 of high spatial resolution observation networks, especially those of ozone precursors. Spatially, 773 774 the addition of ozone monitoring sites along the coastline of LI would benefit the understanding of the spatial extent of ozone pollution and the land/sea air interactions. As inferred from the 775 776 LISTOS NO₂ and HCHO column retrievals, the current ozone monitoring network may have potentially overlooked the high ozone plume on the north shore of LI, where extreme high ozone 777 778 levels are also likely in addition to the two regions studied above. Extensive field campaigns such as LISTOS are beneficial in studying atmospheric physics and chemistry in coastal urban 779 780 environments and validating/developing 3-D numerical models for civil, scientific, and policy-781 relevant use whenever/wherever observations are not feasible.

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809

810 **Open Research**

811 Key scripts to cluster meteorological conditions and corresponding cluster results, as well as supporting materials, are openly accessible under the MIT License 812 other at https://doi.org/10.5281/zenodo.10668892 (Luo & Lu, 2024). 813

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958	

Figure 1.



Figure 2.



Figure 3.



Figure 4.

(a) Wind TS

(b) Wind roses



Figure 5.



(d) Moderate SB, 54 days



(e) Moderate O, 27 days





30



40 45 50 Ozone (

35



(f) Moderate S, 25 days





(g) Moderate W, 26 days





Ozone (ppb)

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55

65

Figure 6.



Figure 7.

(a) Hot SB 07:00

\rightarrow 10 m/s (b) Hot W_sb 07:00 \rightarrow 10 m/s





→ 10 m/s (c) Hot S 07:00



Figure 8.

(a) Hot SB DMA8



(b) Hot W_sb DMA8

(e) Hot W_sb Pex



(d) Hot SB Pex













Figure 9.



Figure 10.



(d) Hot W_sb 11:00



→ 4 m/s

(g) Hot S 11:00 \rightarrow 4 m/s Min Med Max -1.0 3.1 13.3



Figure11.

Figure 12.

Figure 13.

(a) 08:23-10:23 NO₂

(b) 08:23-10:23 HCHO

(e) 12:00-14:11 HCHO

(d) 12:00-14:11 NO₂

(c) 08:23-10:23 HCHO/NO₂

(f) 12:00-14:11 HCHO/NO₂

Figure 14.

Figure 15.

(a) 07:08-09:59 NO₂

(b) 07:08-09:59 HCHO

(d) 11:47-14:41 NO₂

(c) 07:08-09:59 HCHO/NO₂

(f) 11:47-14:41 HCHO/NO₂

